Retrieval of Cloud Thermodynamic Phase Using IR Channels

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Jean-Marie Lalande Emmanuel Fontaine

Météo France/CNRM/CEMS/Nuages

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Motivations : Cloud Phase determination

State

- liquid water
- ice water
- supercooled water (liquid water below the freezing point)
- a mix between liquid water and ice water

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- remote sensing applications

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Status

- based on brightness temperature difference $(BT_{\sim 8.7} BT_{\sim 10.8})$
- investigation for the new FCI radiometer on MTG is necessary
- new approach based on spectral variability of cloud absorption
- study based on simulated radiance with RTTOV

Background on retrieving cloud thermodynamic phase

Absorption of electromagnetic radiation depends on $Im(n_r)$



- $\lambda \in [8.5, 10] \ \mu m$: liquid water and ice absorb approximately equally
- $\lambda \in [11, 12.5] \ \mu m$: ice absorbs more strongly than water.
- When everything else being equal (atmosphere, effective radius, ...)

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Principle of cloud phase retrieval

$$\left[\mathrm{BT}_{\sim 8.7} - \mathrm{BT}_{\sim 10.8}\right]^{\textit{lce}} > \mathrm{BTD}_{\mathrm{thresh}} > \left[\mathrm{BT}_{\sim 8.7} - \mathrm{BT}_{\sim 10.8}\right]^{\textit{Liquid}}$$

- Thresholds depend on satellite, time of the day, geolocalization, atmosphere ...
- Thresholds are set up by experience (cf. ATBD)
- Good enough for MSG-Seviri and other GEO (GOES-16, Himawari)
- What about MTG-FCI?

A simulation study with the ECMWF diverse profile data

Simulation study

- Diverse profile datasets ECMWF 137 levels: 25000 profiles
- We use 4348 of realistic 1-dimensional clouds



- Super fast radiative transfer code RTTOV (absorption/emission + diffusion by clouds)
- Ice cloud: parameterization Baum + Boudala (heritage from H. Le Gléau)
- GEO sensors: MSG, GOES-16, Himawari, MTG



- The MSG 10.8 μm band has been shifted towards lower wavelength on Himawari and MTG
- MTG is missing the 11 $\mu{\rm m}$ band



 $\begin{array}{l} \textit{MSG} \ \textit{/ vza=0} \\ \lambda_1 = 8.7, \ \lambda_2 = 10.8 \end{array}$

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The β -ratio approach: Theory

Relationship between cloud emissivity and cloud effective extinction coefficient (Inoue, *JMSJ*, 1984):

$$\epsilon(\lambda) = 1 - \exp(-\beta(\lambda)z)$$

Equating between 2 wavelengths (spectral bands) λ_1 and λ_2 leads to:

$$\beta' = \frac{\beta(\lambda_1)}{\beta(\lambda_2)} = \frac{\ln\left[1 - \epsilon(\lambda_1)\right]}{\ln\left[1 - \epsilon(\lambda_2)\right]}$$

From the Shwarzschild equation we can solve for the effective cloud emissivity ϵ :

$$\epsilon(\lambda) = \frac{\mathcal{R}_{\rm obs}(\lambda) - \mathcal{R}_{\rm clr}(\lambda)}{\underbrace{\left[\mathcal{T}_{\rm ac}(\lambda)\mathcal{B}(\lambda, \mathcal{T}_{\rm eff}) + \mathcal{R}_{\rm ac}(\lambda)\right]}_{\mathcal{R}_{\rm blk}(\lambda)} - \mathcal{R}_{\rm clr}(\lambda)}$$

- RTTOV-DOM (Discrete Ordinate Method for diffusion)
- \mathcal{R}_{clr} : Clear sky (absorption/emission)
- $\mathcal{R}_{\rm blk}:$ Blackbody radiation transmitted to TOA + Above cloud emission



The β -ratio computation

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- Computed from NWP output (Temperature, Pressure, gas concentration)
- β-ratio is more sensitive to cloud phase (removal of clear sky effect)



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- Classification algorithm: Logistic Regression/SVM using multiple predictors



So far...

- Simulation study on all GEO sensors using a 1D "realistic" clouds
- Limitation of the traditional Brightness temperature difference
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- 4 MTG data
 - When available . . .
 - Probably more tuning required
 - Final algorithm . . .